Variability in Mesoscale Physical Activity in the Northern California Current and its Effects on Biological Distributions



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Introduction

Mesoscale circulation features contribute significantly to cross-shelf transport of biological production from nearshore areas to the deep sea. As part of the U.S. GLOBEC Northeast Pacific program, we are studying the linkage between mesoscale features and the distribution of chlorophyll and secondary production in upwelling areas of Oregon and northern California. Dynamic mesoscale physical features develop through the summer upwelling season in coastal areas of the California Current System and move offshore. The larger of these are generated in fairly predictable, topographicallycontrolled locations and persist from weeks to months.

In our study area, mesoscale activity and zooplankton biomass and distribution vary seasonally and interannually, with peak biomass typically coinciding with peak mesoscale physical activity (in late summer). Variation in the seasonal timing and magnitude of mesoscale activity can therefore affect cross-shelf transfer of biomass and hence can affect the overall productivity of the California Current System. Our goal in this presentation is to characterize the spatial and temporal variability in mesoscale circulation to further our understanding of the dominant mechanisms controlling biological distributions.



(Figure 1): August 2000: An example of the relationship between mesoscale circulation features and biological fields. Panels are (left to right) copepod biomas from net-tows during a U.S. GLOBEC cruise, chlorophyll a from SeaWiFS, sea surface temperature from the AVHRR, and sea surface height from AVISO.



Sea surface height (SSH) data and estimates of gridding error (Figure 2) come from AVISO (Archiving, Validation and Interpretation of Satellite Oceanographic data) <u>http://www.aviso.oceanobs.com</u>):

- Weekly fields of gridded sea surface anomalies
- Oct 1992 to Dec 2004
- TOPEX/ERS/ENVISAT/and JASON altimeters (as available)
 SSH gridded over 6 weeks of data, weighted to the central date

Specific locations for wavelet analysis were chosen by examining the average error in the gridded fields (Figure 2) and selecting locations with low temporally-averaged errors that were also wellspaced meridionally and zonally at similar distances from shore.

Figure 2. Gridding error in AVISO sea surface height anomaly fields averaged over the entire time series – expressed as % of total variance. X = locations chosen for wavelet analyses. Note that error is lowest along TOPEX ground tracks.

<u>Wavelet analysis</u>: Wavelet methods were based on Torrence and Compo (1998). The wavelet power spectrum is constructed by convolving the time series (X_i) with a scaled wavelet (ψ_i), to produce a matrix of N data points by *s* scale factors:

$$W_n(s) = \sum_{n'=0}^{N-1} x_n \psi'^* \left[\frac{(n'-n)\partial t}{s} \right] \text{ and the power is defined as: } \left| W_n(s) \right|^2$$

Here, we use the Morlet wavelet (with ω_0 =6), which is comprised of a sine wave modified by a Gaussian: $\psi_0(\eta) = \pi^{-1/4} e^{i\omega_0 \eta} e^{-\eta^2/2}$

The local wavelet power shown in figures is the square of the wavelet coefficients normalized by the variance of each time series. Significance was tested at each location by comparing the wavelet variance to a red-noise background spectrum defined by the variance and length of each individual time series.

